



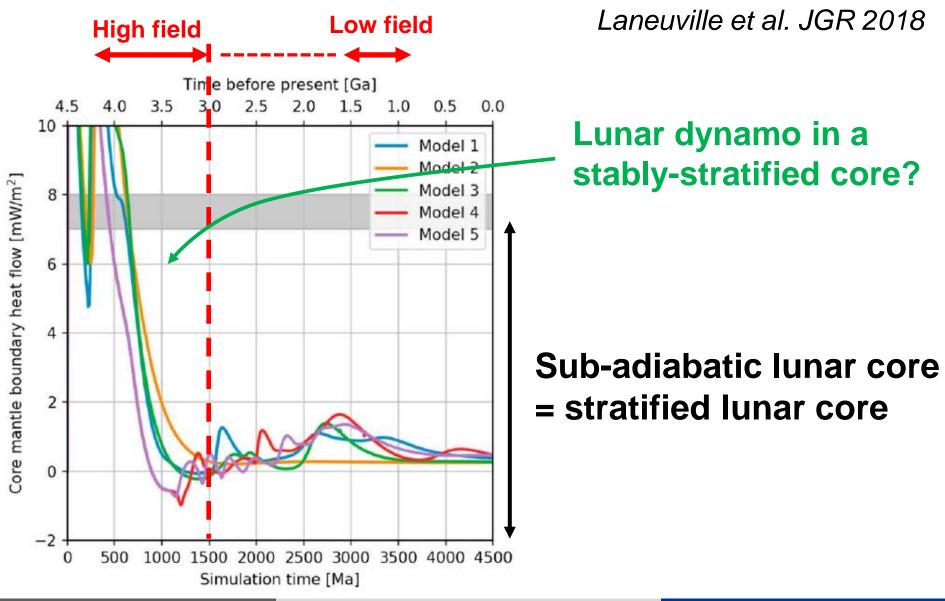
Rotating convection in stably-stratified planetary cores

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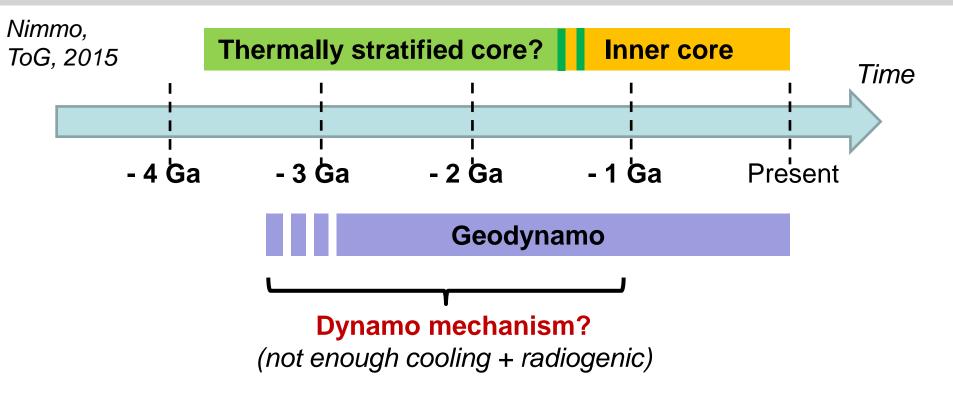
The Core of the Moon, Marseille, May 20-22, 2019

A thermally stratified lunar core?



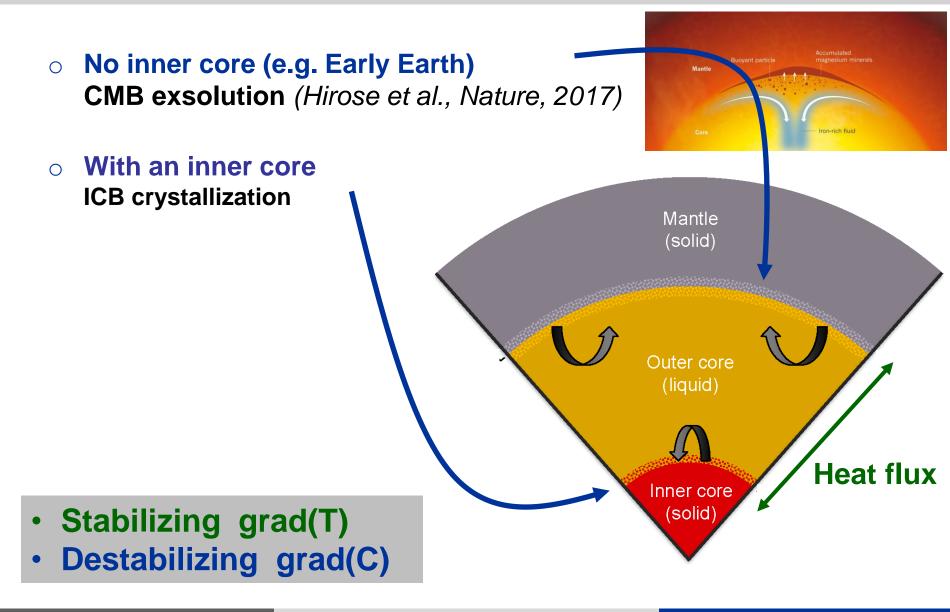
David Cébron

A thermally stratified Earth's core?

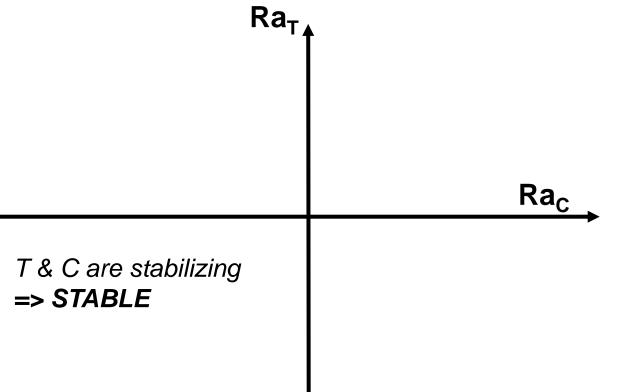


 Rourke et al. (2017): "how to power convection in the core and thus a dynamo for the vast majority of the Earth's history remains one of the most pressing puzzles in geophysics"

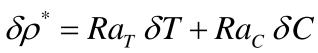
Two sources of buoyancy in planetary liquid cores



Dimensionless buoyancy



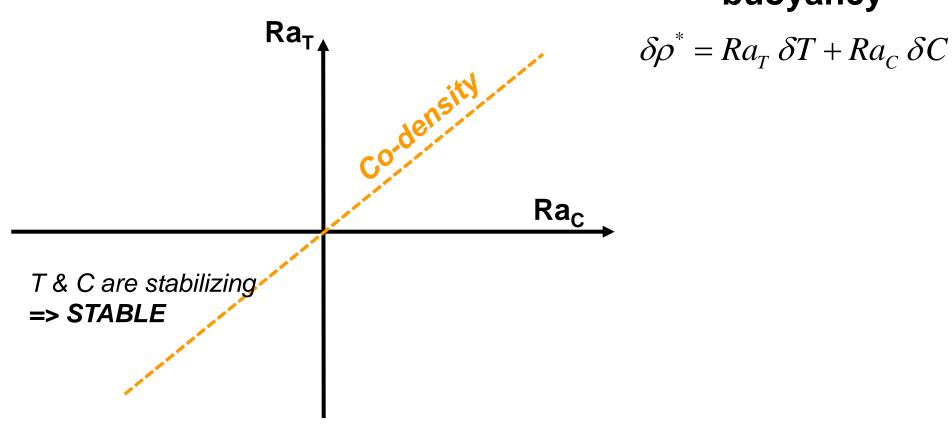
buoyancy



Regimes & Onset

Usual dynamo simulations: co-density

Dimensionless buoyancy



Regimes & Onset

- Usual dynamo simulations: co-density
- Less usual dynamo simulations

Ra_T T & C are destabilizing Ra_c T & C are stabilizing => STABLE

Dimensionless buoyancy

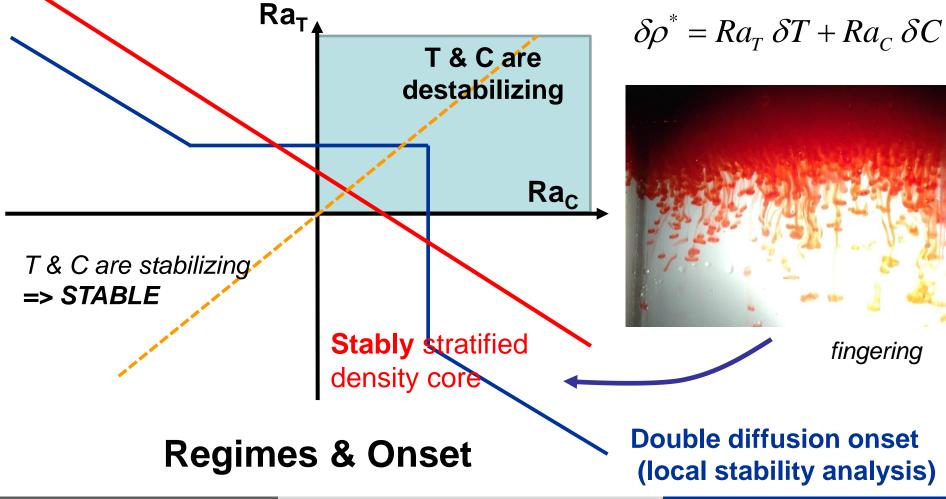
$$\delta \rho^* = Ra_T \,\delta T + Ra_C \,\delta C$$

Regimes & Onset

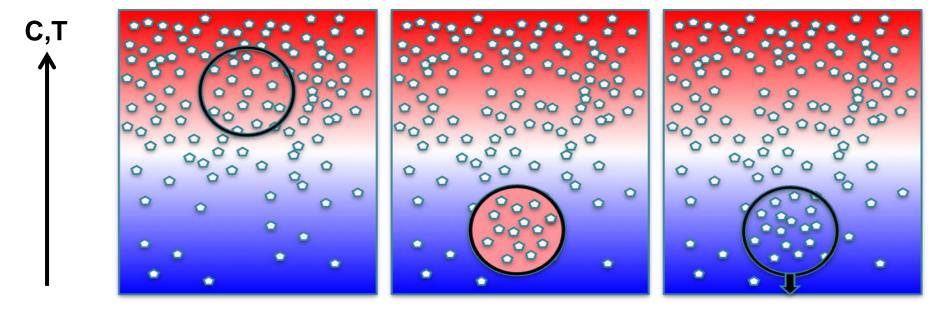
Usual dynamo simulations: co-density Dimensionless Less usual dynamo simulations buoyancy Ra_τ $\delta \rho^* = Ra_T \,\delta T + Ra_C \,\delta C$ T & C are destabilizing Ra_c No flow? T & C are stabilizing => STABLE **Stably** stratified density core **Regimes & Onset**

Usual approach: co-density

Dimensionless buoyancy



A little known double-diffusion in the Earth's core



Fast diffusion of T wrt. C

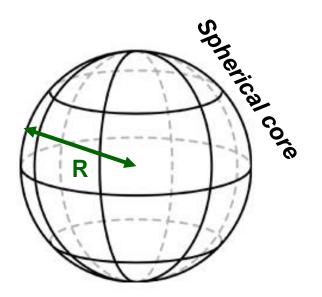
Instability well known in ocean & stellar physics

But almost no study in a **rotating spherical geometry** (*Manglik*+10, *Net*+12, *Bouffard's PhD*)

Double diffusion in the Early Earth's core?

• **Scales:** R, R²/v, T, C, such that
$$T_0(r) = \frac{1 - r^2}{Pr}$$
, $C_0(r) = \frac{1 - r^2}{Sc}$

$$\begin{aligned} \frac{\partial \boldsymbol{u}}{\partial t} + (\boldsymbol{u} \cdot \boldsymbol{\nabla})\boldsymbol{u} &= -\frac{2}{Ek} \mathbf{1}_{z} \times \boldsymbol{u} - \nabla p + \boldsymbol{\nabla}^{2} \boldsymbol{u} \\ + (Ra_{T} \Theta + Ra_{C} \xi) \ r \mathbf{1}_{r}, \\ \frac{\partial \Theta}{\partial t} + (\boldsymbol{u} \cdot \boldsymbol{\nabla})\Theta &= \frac{1}{Pr} \left(2 \ \boldsymbol{r} \cdot \boldsymbol{u} + \nabla^{2} \Theta \right), \\ \frac{\partial \xi}{\partial t} + (\boldsymbol{u} \cdot \boldsymbol{\nabla})\xi &= \frac{1}{Sc} \left(2 \ \boldsymbol{r} \cdot \boldsymbol{u} + \nabla^{2} \xi \right), \\ \boldsymbol{\nabla} \cdot \boldsymbol{u} &= 0, \end{aligned}$$



$$Ra_{T} = \frac{\alpha_{T}g_{0}Q_{T}R^{6}}{6\nu\kappa_{T}^{2}}, \quad Ra_{C} = \frac{\alpha_{C}g_{0}Q_{C}R^{6}}{6\nu\kappa_{C}^{2}}, \qquad Ek = \frac{\nu}{\Omega_{s}R^{2}}, \quad Pr = \frac{\nu}{\kappa_{T}}, \quad Sc = \frac{\nu}{\kappa_{C}}$$
Planetary cores:
<10⁻¹⁰ L=Sc/Pr>10^{3}

Equations

0

Double diffusion in the Early Earth's core?

• Scales: R, R²/v, T, C, such that
$$T_0(r) = \frac{1 - r^2}{Pr}$$
, $C_0(r) = \frac{1 - r^2}{Sc}$

• Equations

$$\frac{\partial u}{\partial t} + (u \cdot \nabla)u = -\frac{2}{Ek} \mathbf{1}_{z} \times u - \nabla p + \nabla^{2} u \\
+ (Ra_{T} \Theta + Ra_{C} \xi) r \mathbf{1}_{r},$$

$$\frac{\partial \Theta}{\partial t} + (u \cdot \nabla)\Theta = \frac{1}{Pr} (2r \cdot u + \nabla^{2}\Theta),$$

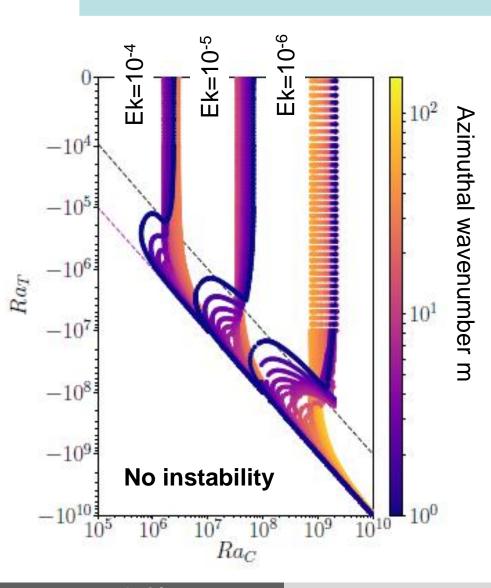
$$\frac{\partial \xi}{\partial t} + (u \cdot \nabla)\xi = \frac{1}{Sc} (2r \cdot u + \nabla^{2}\xi),$$

$$\nabla \cdot u = 0,$$
Summary constraints of the second second

$$Ra_T = \frac{\alpha_T g_0 \mathcal{Q}_T R^6}{6\nu\kappa_T^2}, \quad Ra_C = \frac{\alpha_C g_0 \mathcal{Q}_C R^6}{6\nu\kappa_C^2}, \qquad Ek = \frac{\nu}{\Omega_s R^2}, \quad Pr = \frac{\nu}{\kappa_T}, \quad Sc = \frac{\nu}{\kappa_C}$$

Codes: open-source, pseudo-spectral => SINGE (linear, by J. Vidal)
 & XSHELLS (non-linear, by N. Schaeffer)

Onset in a rotating sphere at L=Sc/Pr=10



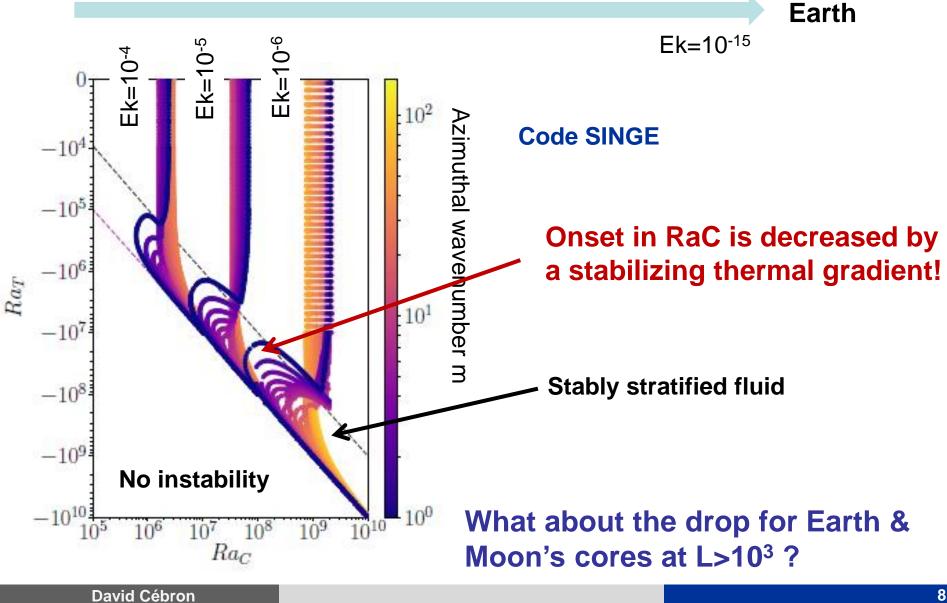
Code SINGE

Monville et al., GJI, subm.

Earth

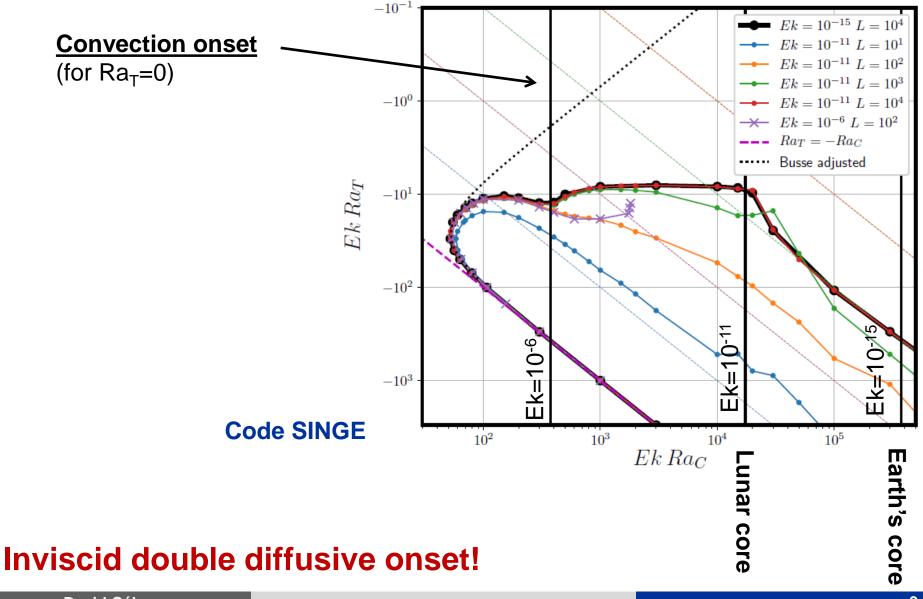
Ek=10⁻¹⁵

Onset in a rotating sphere at L=Sc/Pr=10



Onset in a rotating sphere at core values

L=Sc/Pr



David Cébron

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Inviscid convection domain

L=Sc/Pr

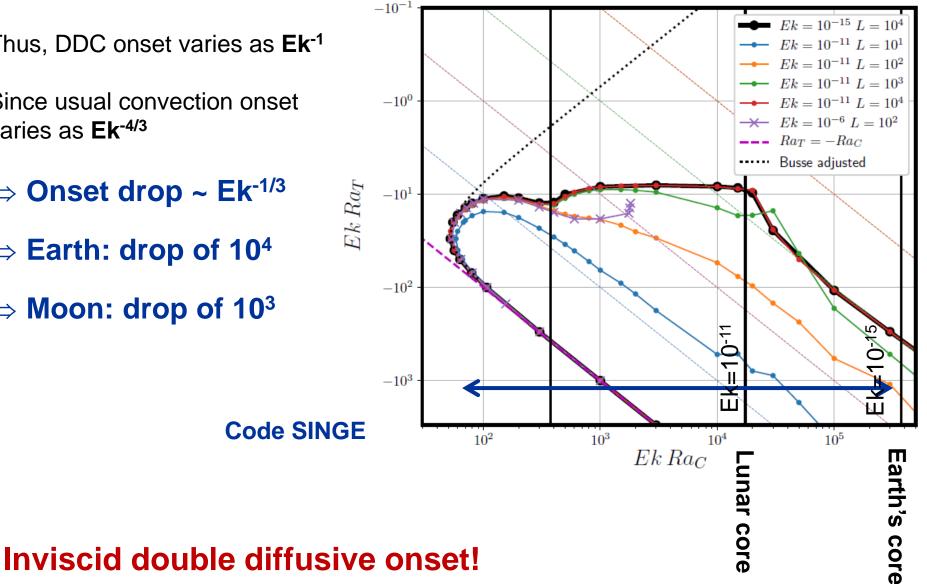
Thus, DDC onset varies as **Ek**⁻¹

Since usual convection onset varies as **Ek**^{-4/3}

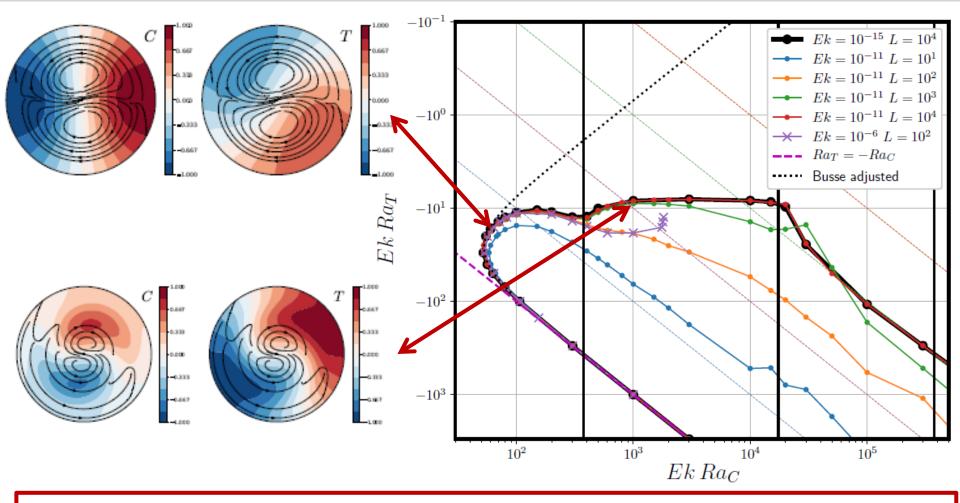
 \Rightarrow Onset drop ~ Ek^{-1/3}

 \Rightarrow Earth: drop of 10⁴

 \Rightarrow Moon: drop of 10³



Eigenmodes at the onset



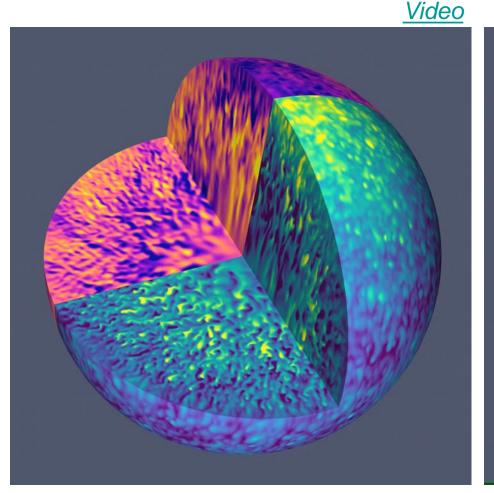
Large-scale inviscid double diffusive flow in the fingering regime

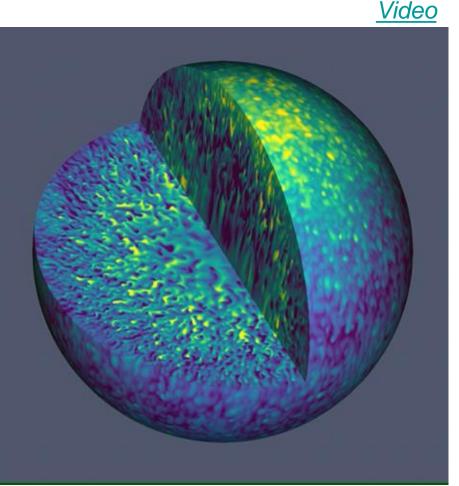
Non-linear regime?

Non-linear rotating double diffusive convection

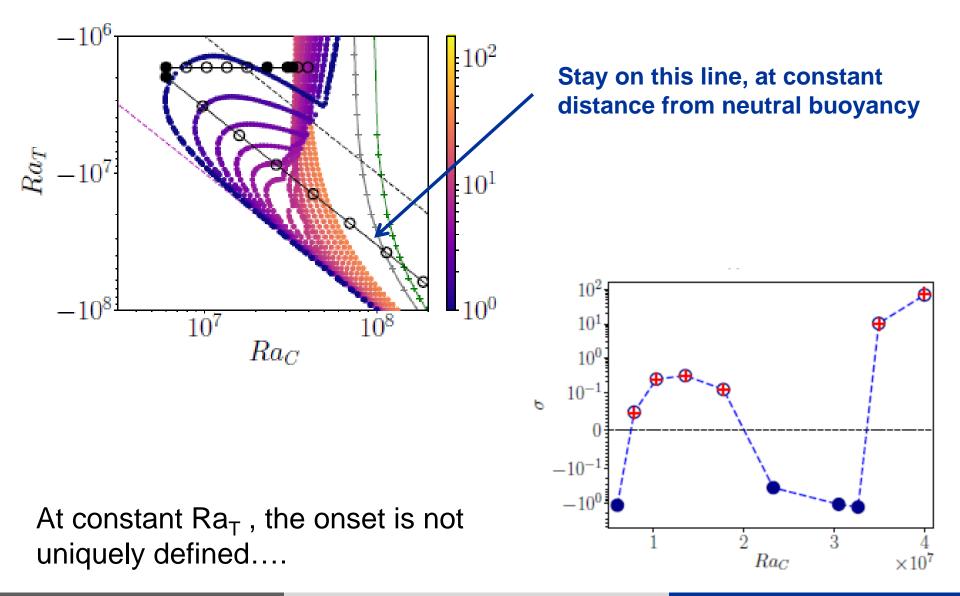
Thermal (upper) & solutal (lower) buoyancy

Vorticity magnitude

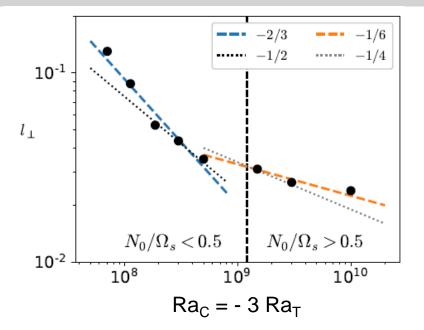




Double diffusive growth rates at L=Sc/Pr=10



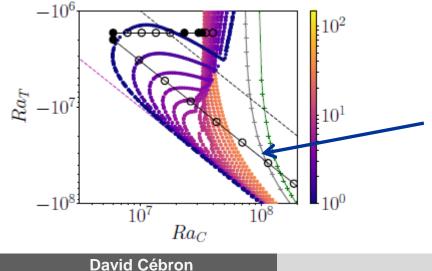
Double diffusive structures & transport at L=Sc/Pr=10



Radko, 2013, non-rotating : -1/4

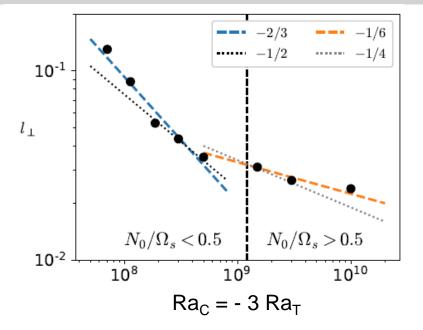
Bouffard, 2017, rotating: -1/2

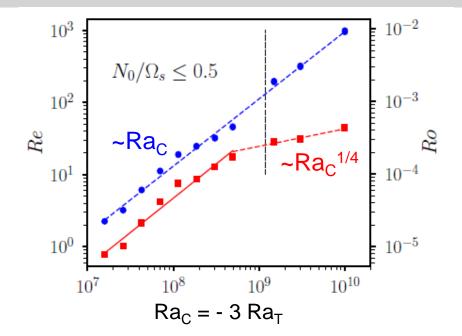
=> Transition between 2 regimes (weakly and strongly stratified)



Stay on this line, at constant distance from neutral buoyancy

Double diffusive structures & transport at L=Sc/Pr=10



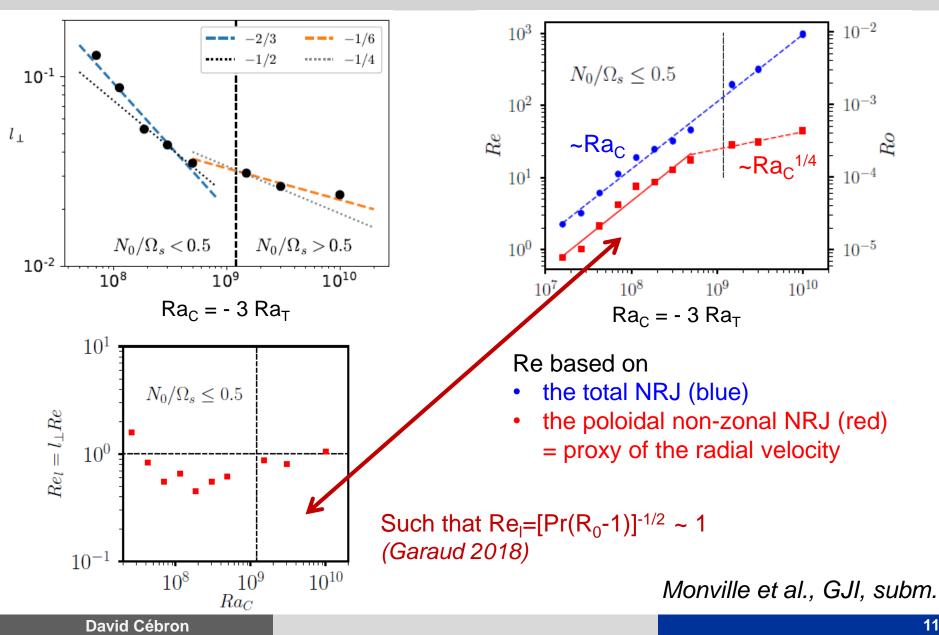


Re based on

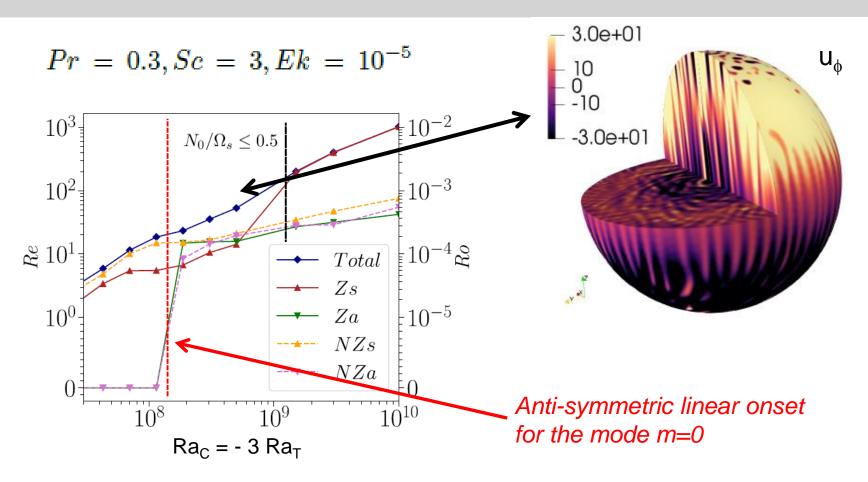
- the total NRJ (blue)
- the poloidal non-zonal NRJ (red)
 = proxy of the radial velocity

Monville et al., GJI, subm.

Double diffusive structures & transport at L=Sc/Pr=10

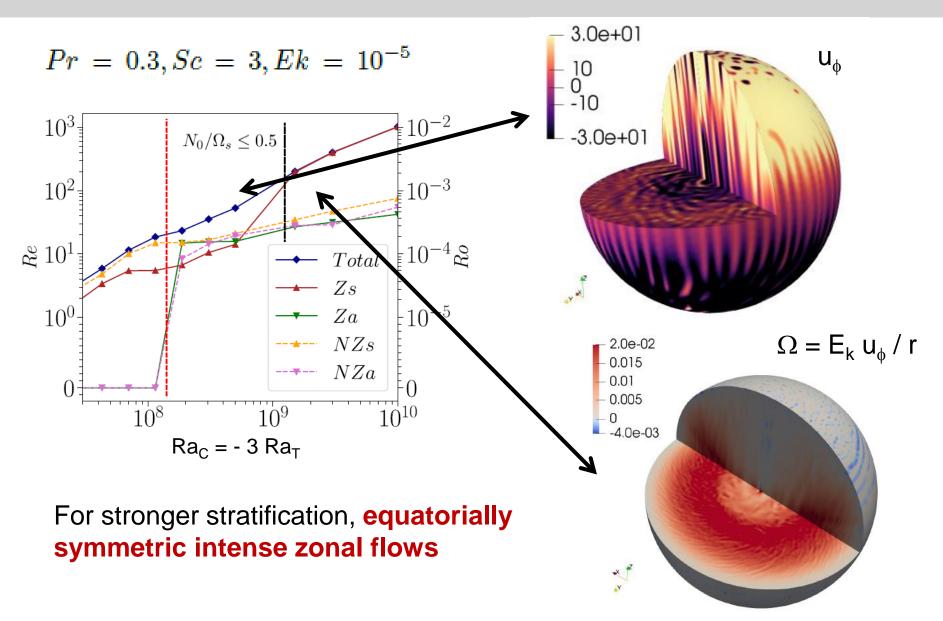


Zonal flows



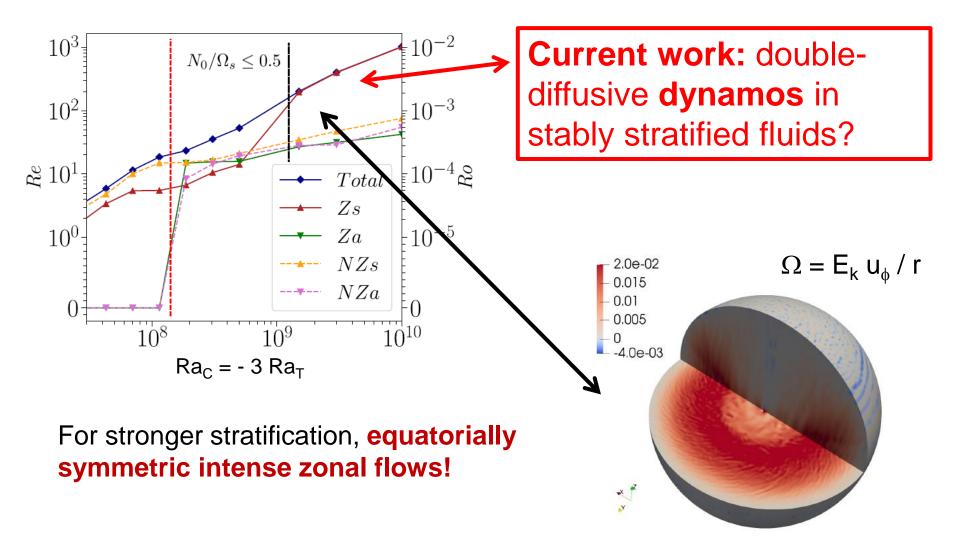
- Far from RDDC onset, equatorially anti-symmetric zonal flow
- Well predicted by a purely linear mechanism

Zonal flows

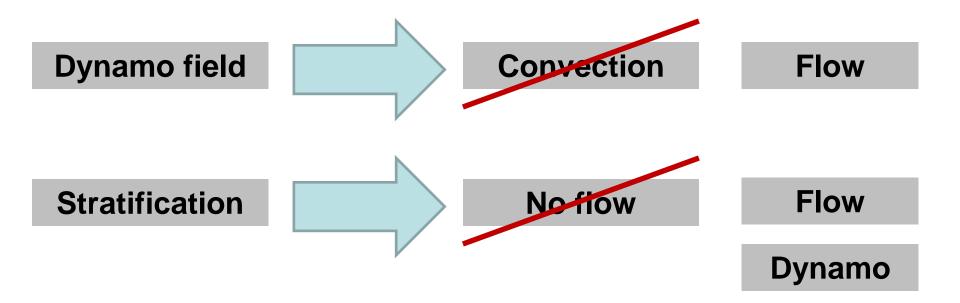


Zonal flows

 $Pr = 0.3, Sc = 3, Ek = 10^{-5}$



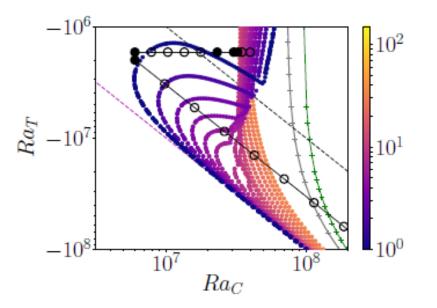
Conclusions



The same for mechanical dynamos...

More details: Monville et al. (2019) arXiv:1902.08523 Code: https://nschaeff.bitbucket.io/xshells/

Double diffusive transport



$$Nu_{T} = \frac{T_{0}(0) - T_{0}(1)}{T_{0}(0) - T_{0}(1) + \Theta_{rms}(0) - \Theta_{rms}(1)},$$

$$Sh = \frac{C_{0}(0) - C_{0}(1)}{C_{0}(0) - C_{0}(1) + \xi_{rms}(0) - \xi_{rms}(1)},$$

